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Ocean-Atmosphere interaction and Interannual monsoon variability

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2nd ACAM Training School: Observation & modeling of atmospheric chemistry & aerosols in the Asian monsoon region

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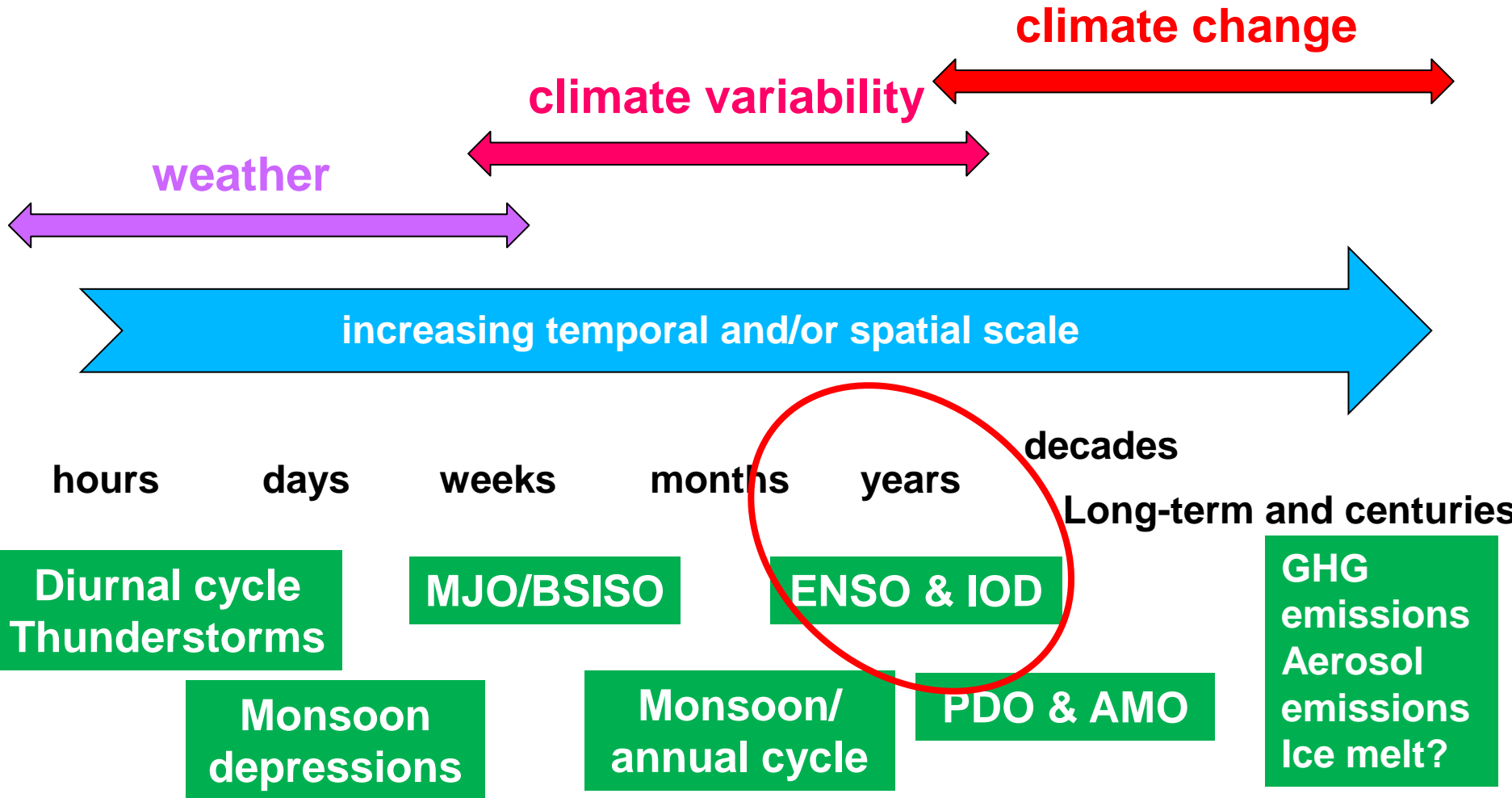


Outline

- 1. Background**
- 2. EASM and ENSO**
- 3. ISM and ENSO**
- 4. Concluding remarks**



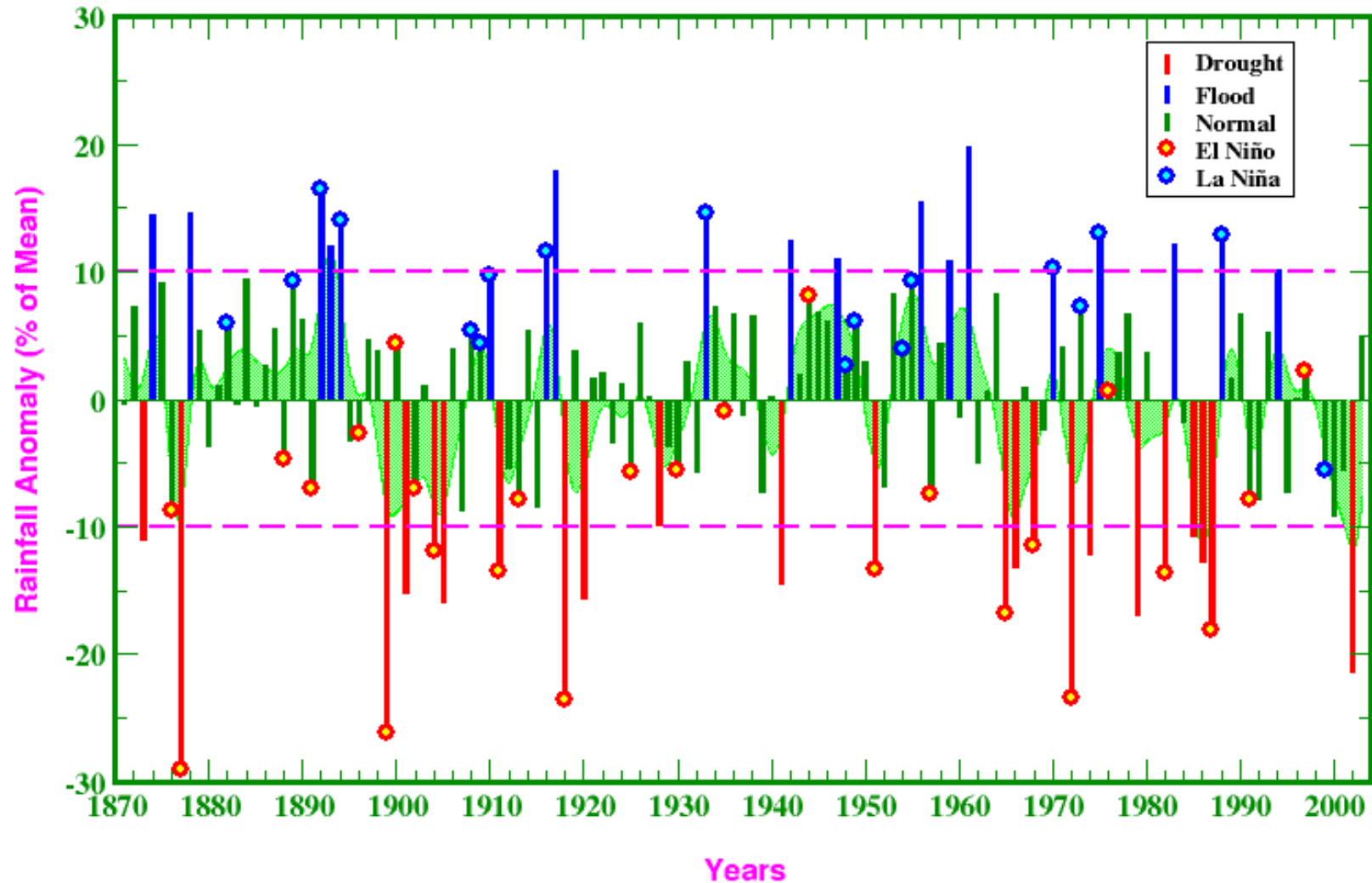
Space and time scales in the monsoon



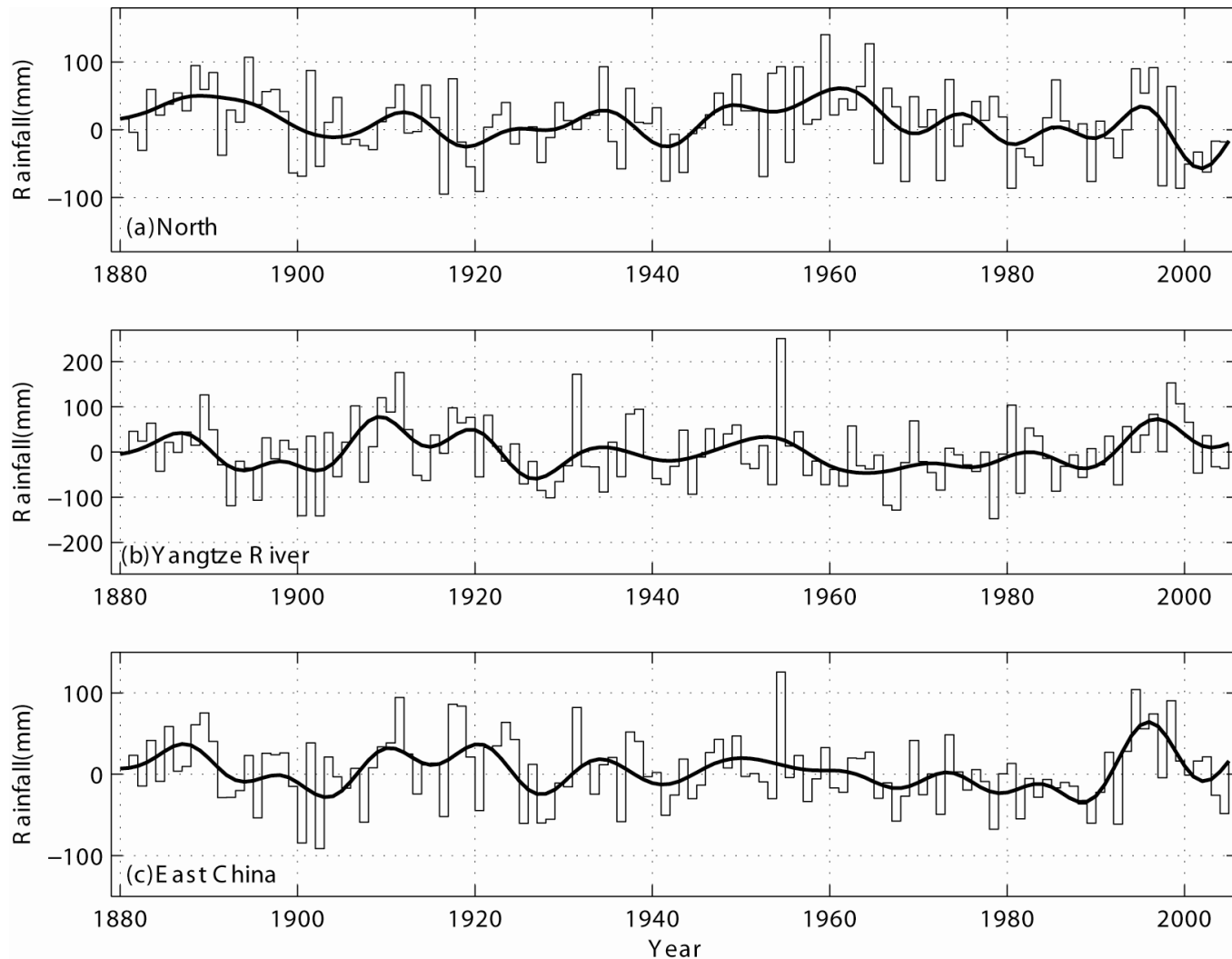
Monsoon-ENSO co-variation

All-India Summer Monsoon Rainfall, 1871-2003

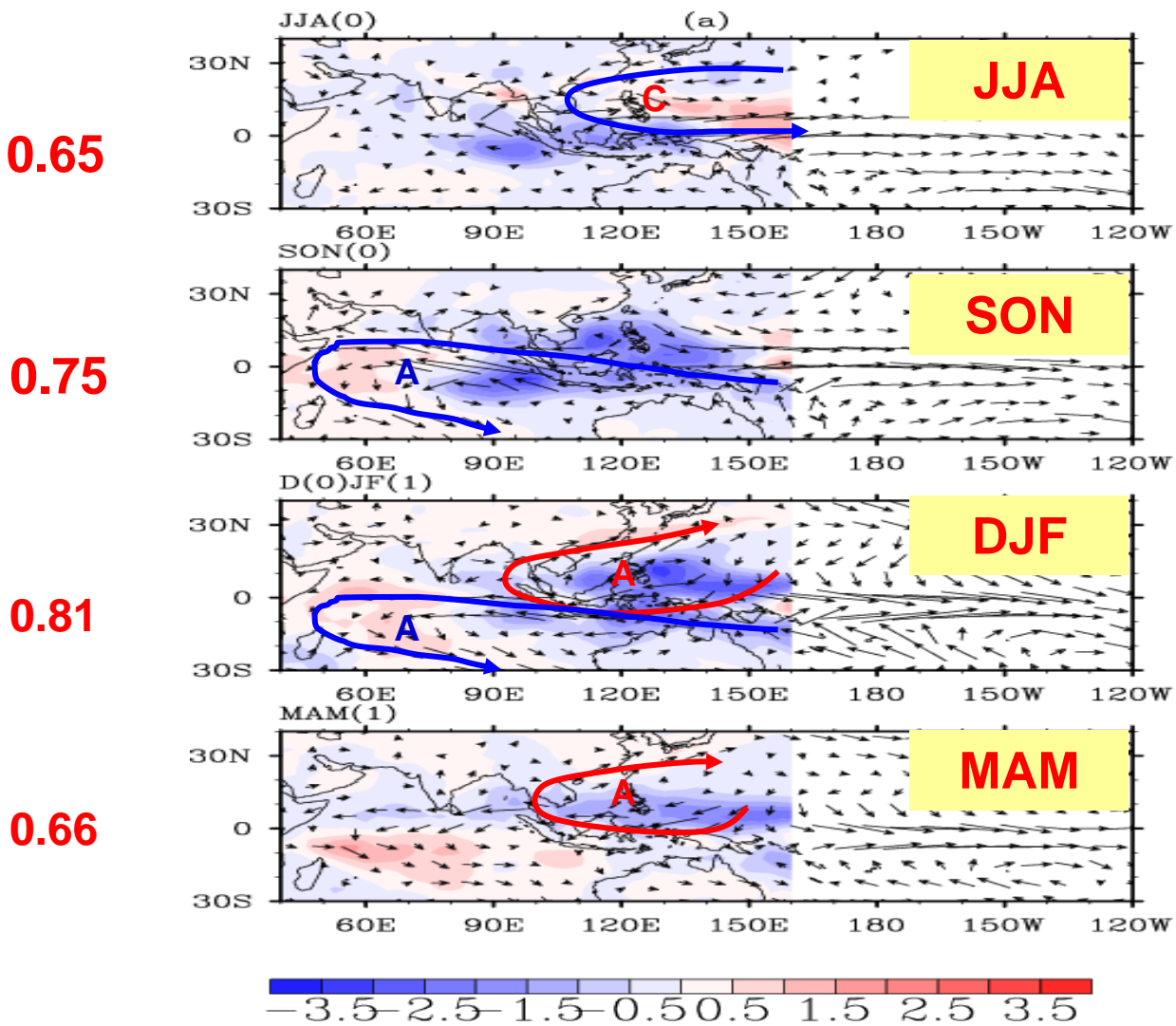
(Based on IITM Homogeneous Indian Monthly Rainfall Data Set)



East Asian summer rainfall



S-EOF modes of MME precipitation and the associated 850hPa wind



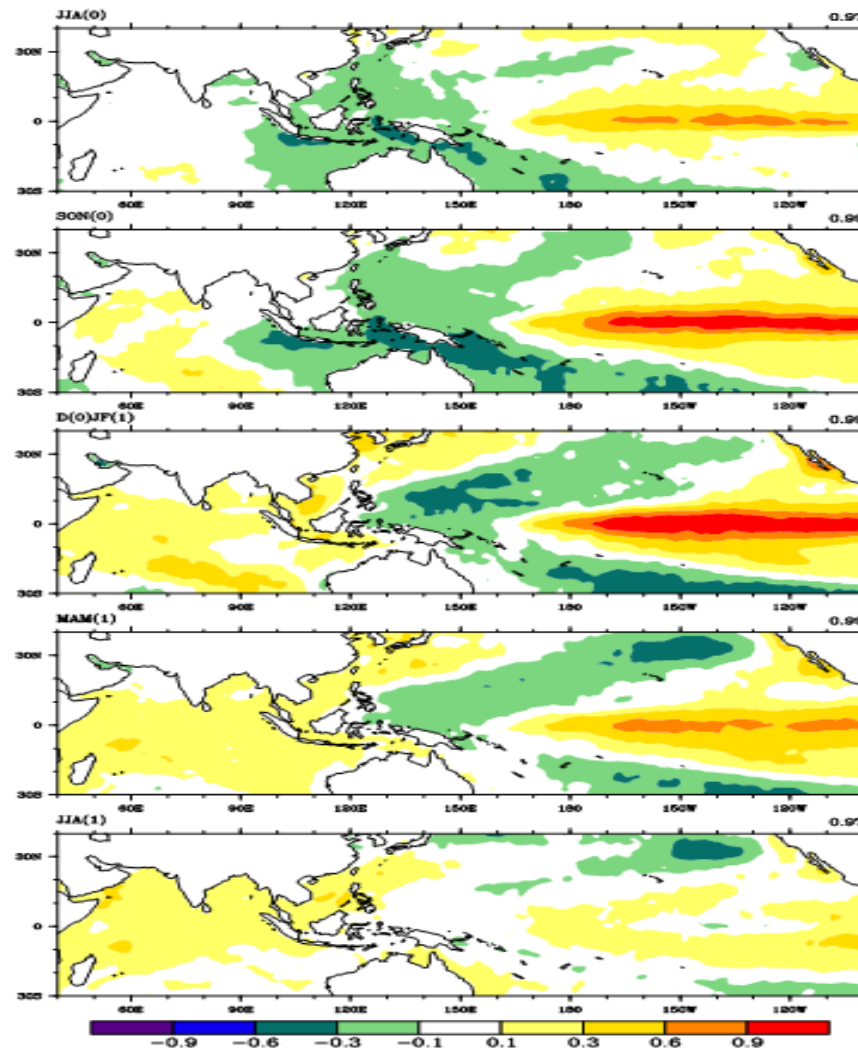
SEOF1(31.3%)

shading: precp

vector: UV850

AMIP MME

SST anomalies in El Nino decaying year



JJA (0)

SON (0)

D(0)JF (1)

MAM (1)

JJA (1)



The key question for interannual monsoon variability is to understand monsoon-ENSO tele-connection

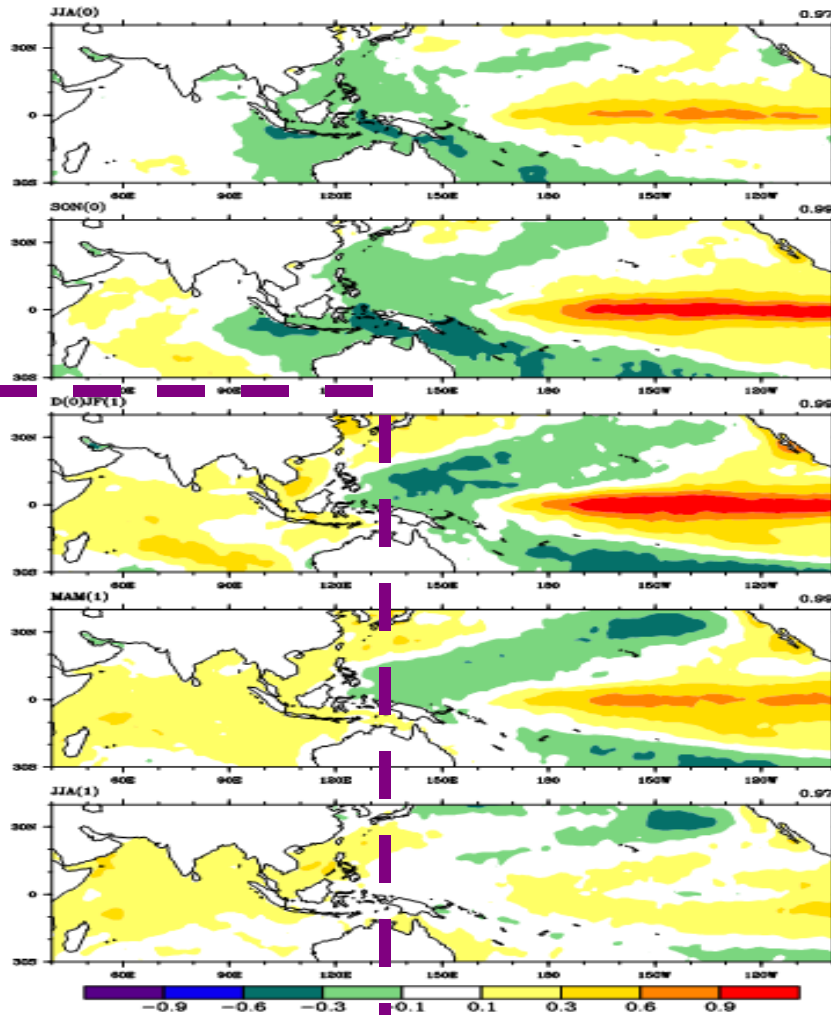


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Indian Ocean SST anomalies versus Pacific SST



DJF

MAM

JJA

The following work highlight Indian Ocean SST forcing to EA climate:

Hu (1997 JGR) ;

Guo Yufu (2004 AAS) ;

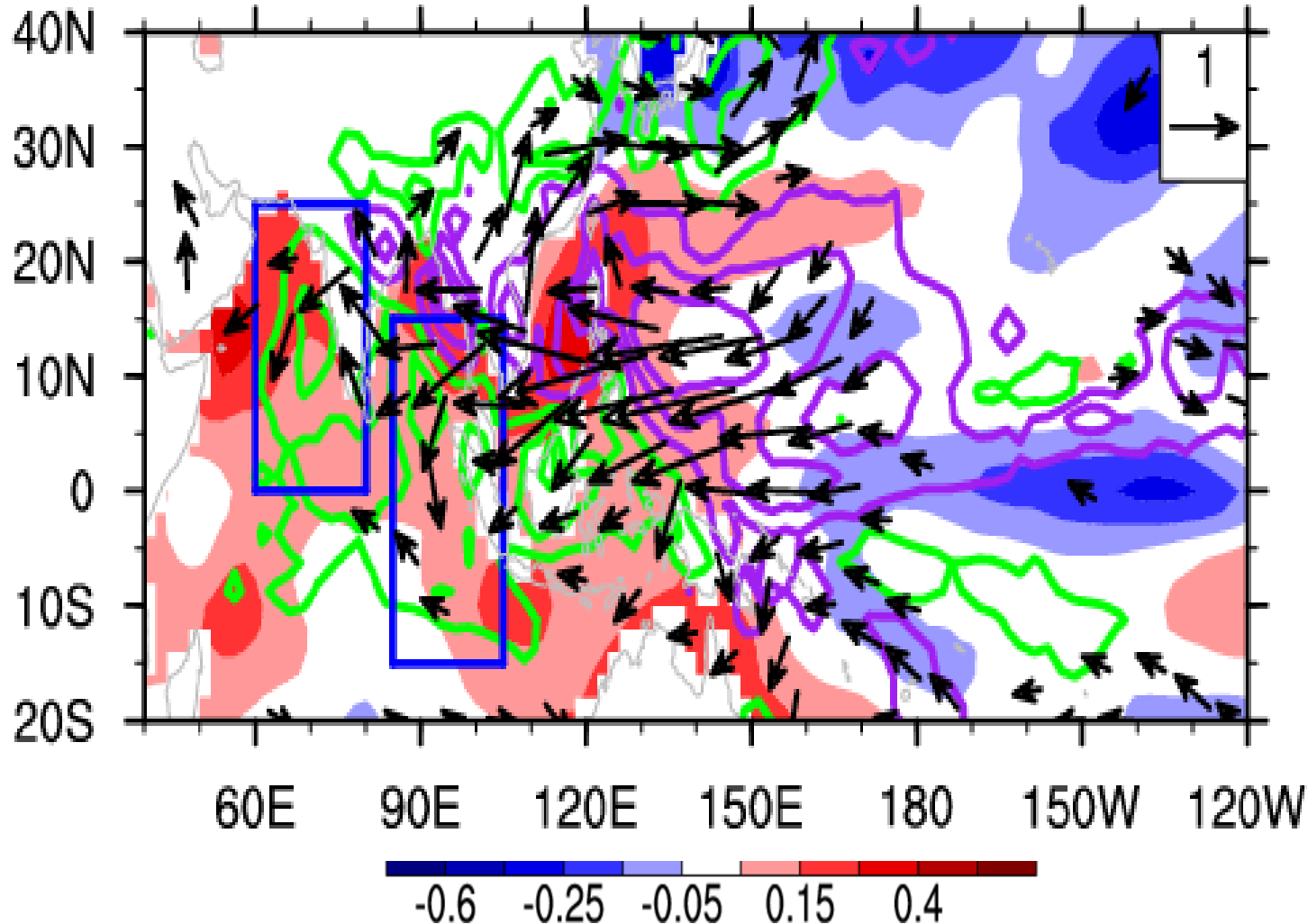
Watanabe & Jin (2007 GRL) ;

Yang et al. (2007)、 Xie et al. (2009) Indian Ocean capacitor;

Wu et al. (2009) seasonal dependence of Indian and western Pacific SSTA

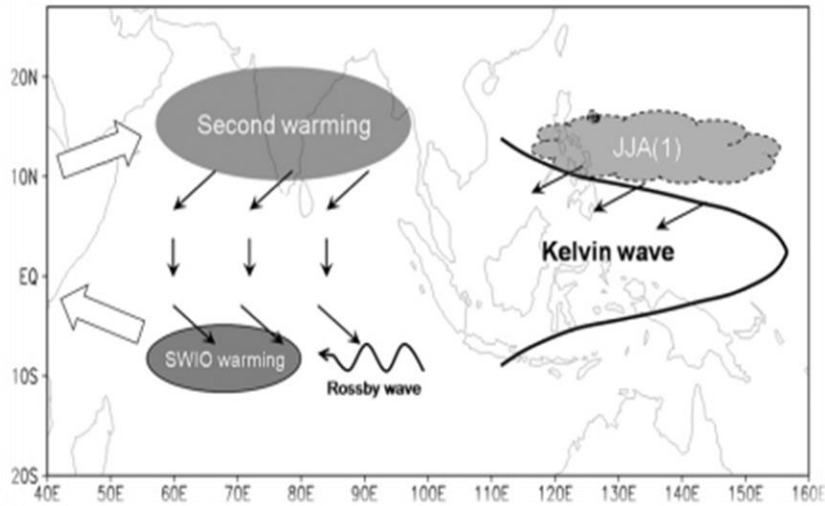


SST (shading), UV850 hPa (Vector), precipitation (contour)



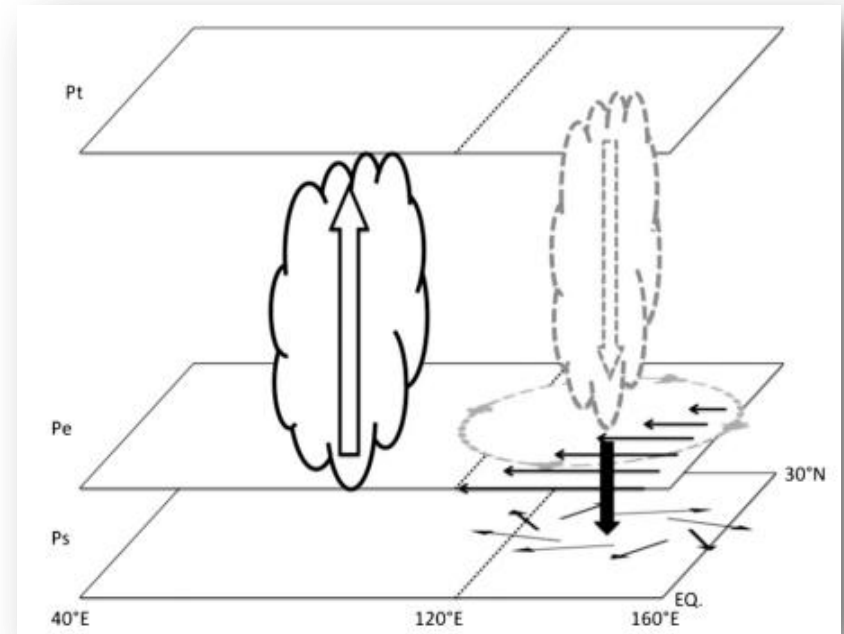


Indian Ocean capacitor effect



(Xie et al. 2009 JC)

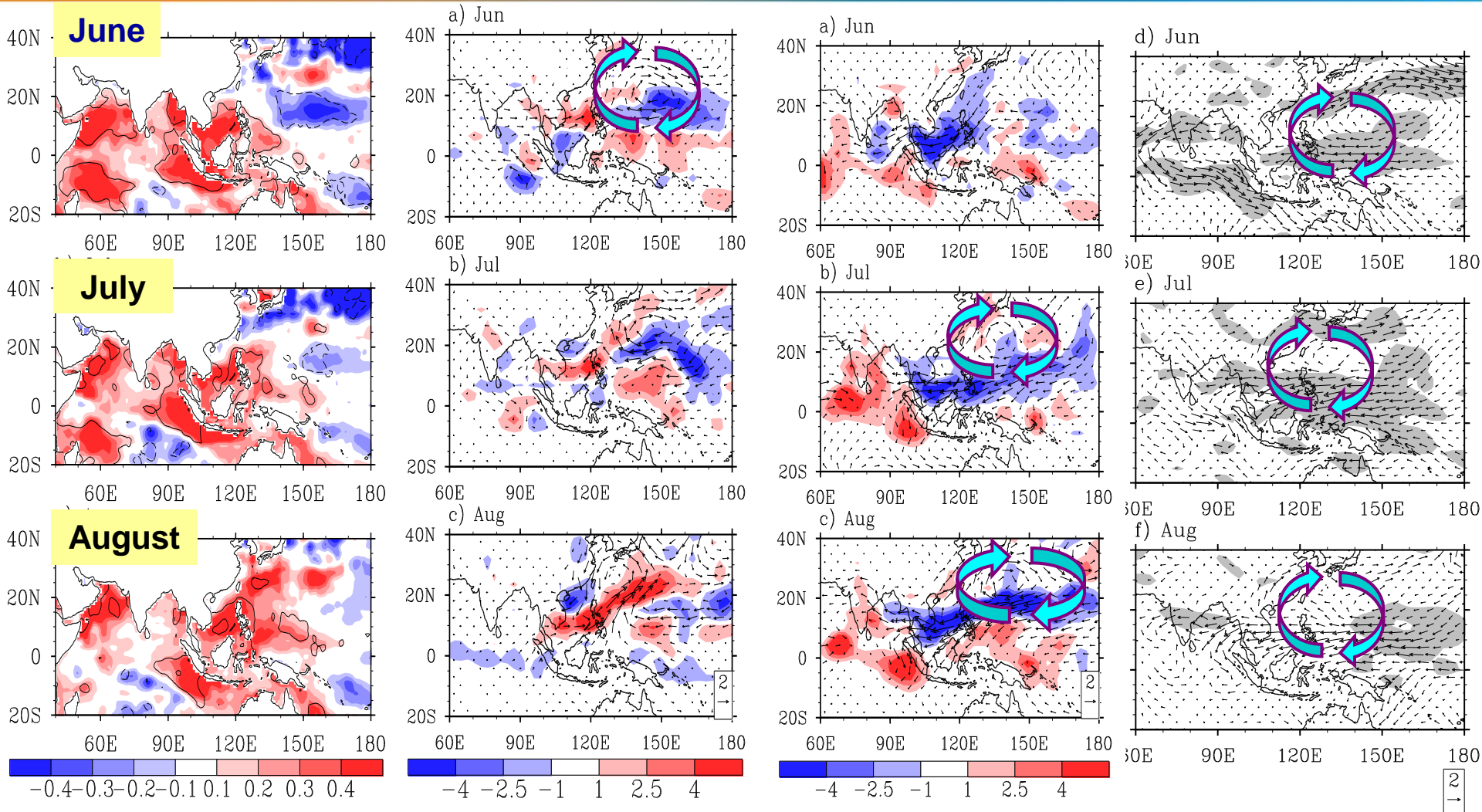
Ekman pumping



(Wu et al. 2009 JC)



Contributions of Indian Ocean and western Pacific SSTA



**SST in El Niño
decaying year**

W. Pac forcing

Indian Ocean forcing

Obs

Point # 1



- ◆ the WNPAC is maintained by the combined effects of the local forcing of the negative SSTA in the WNP and the remote forcing from the IOBM.
- ◆ The former (latter) contribution gradually weakens (enhances) from June to August. The negative SSTA in the WNP is crucial for the maintenance of the WNPAC in early summer.
- ◆ The IOBM plays a crucial role in late summer via the Kelvin wave induced anticyclonic shear and boundary layer divergence.

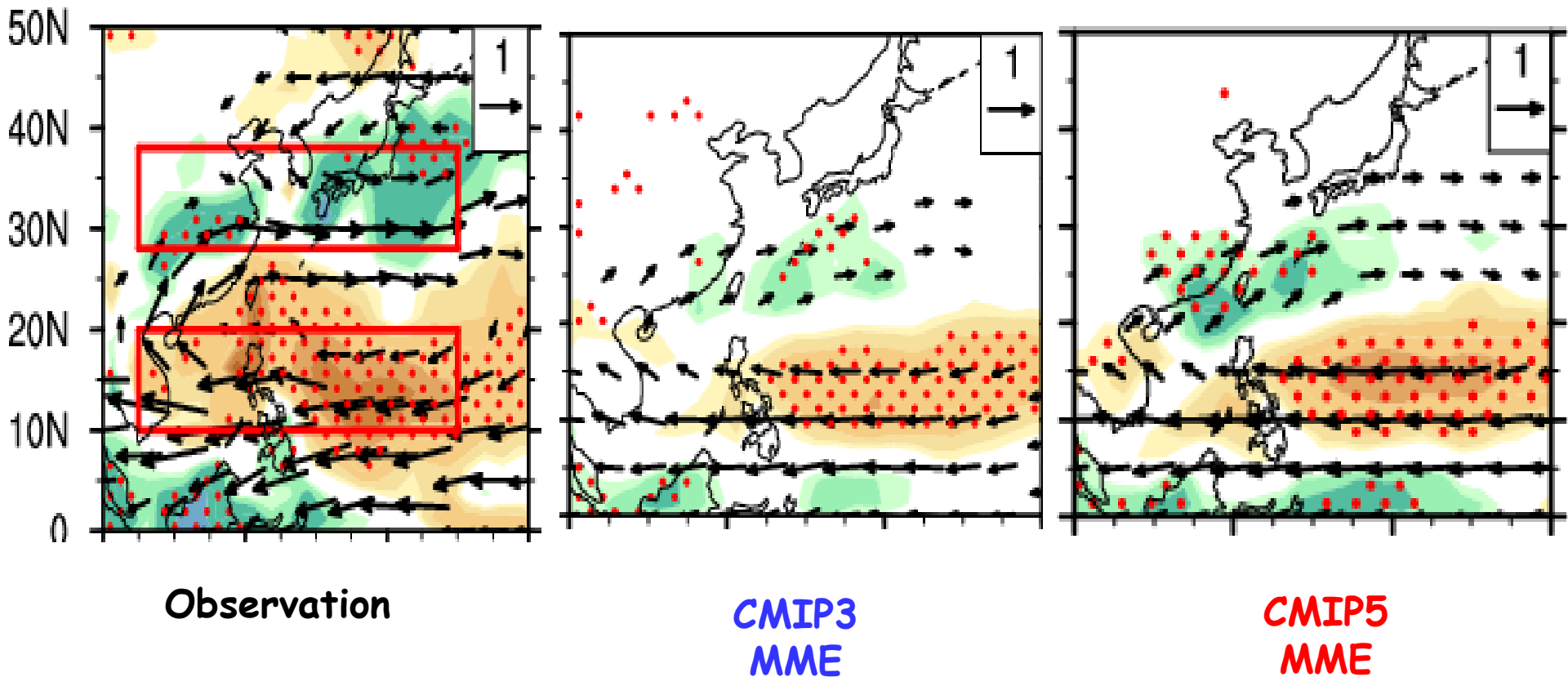


How about climate models' performances?



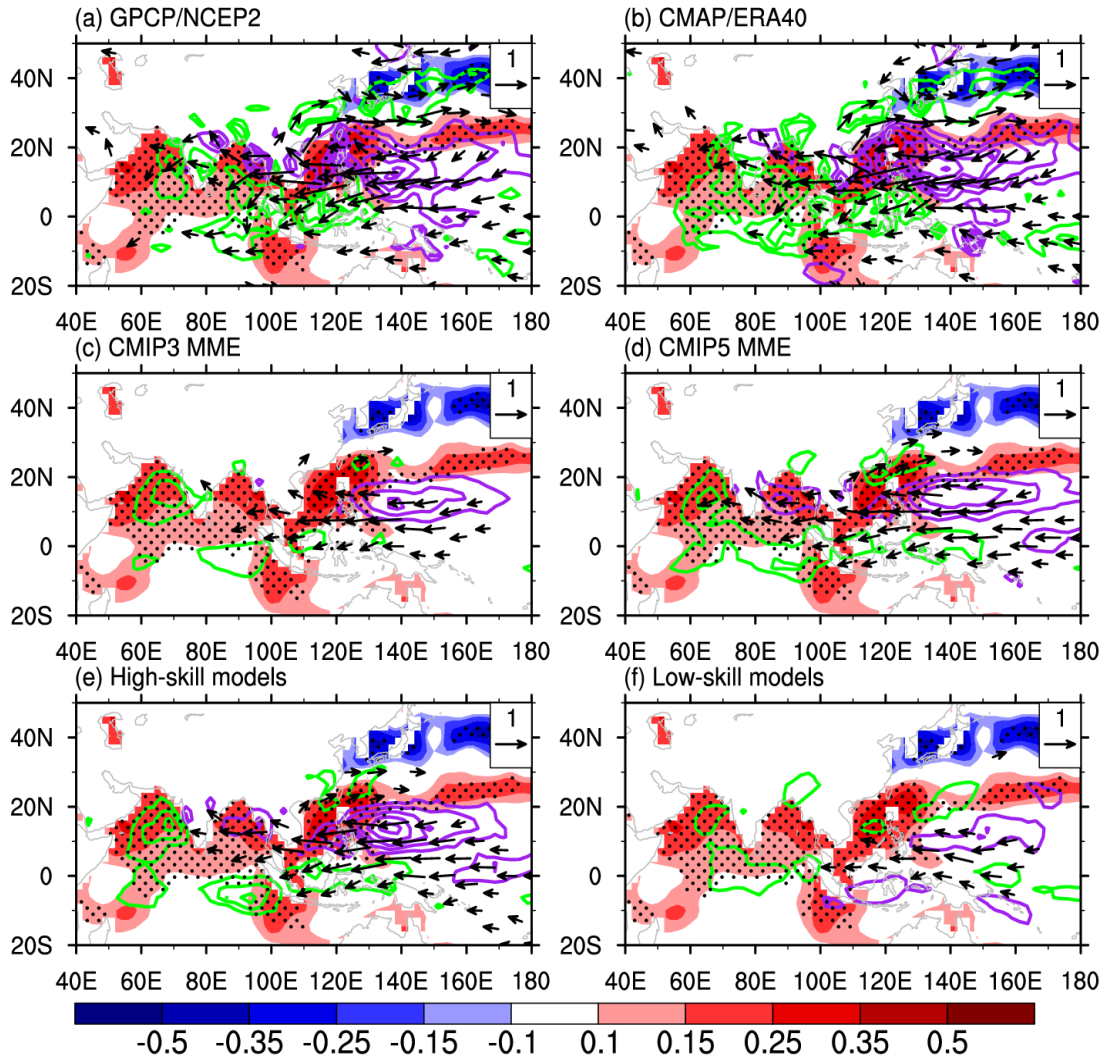


- **13 CMIP3** and **19 CMIP5 AMIP experiments**.
- Observational and reanalysis data:
 - NCEP2: 850 hPa wind, air temperature;
 - GPCP: precipitation;
 - ERSST: SST;
- **Period: 1980 to 1997.**
- All the datasets are interpolated onto common grid
 $2.5^{\circ} \times 2.5^{\circ}$



- **Southward shifts of the W. Pacific Anticyclone and the associated rainfall anomalies over EA; Similar bias in CMIP3 & CMIP5 models**

Indian Ocean-western Pacific anticyclone tele-connection



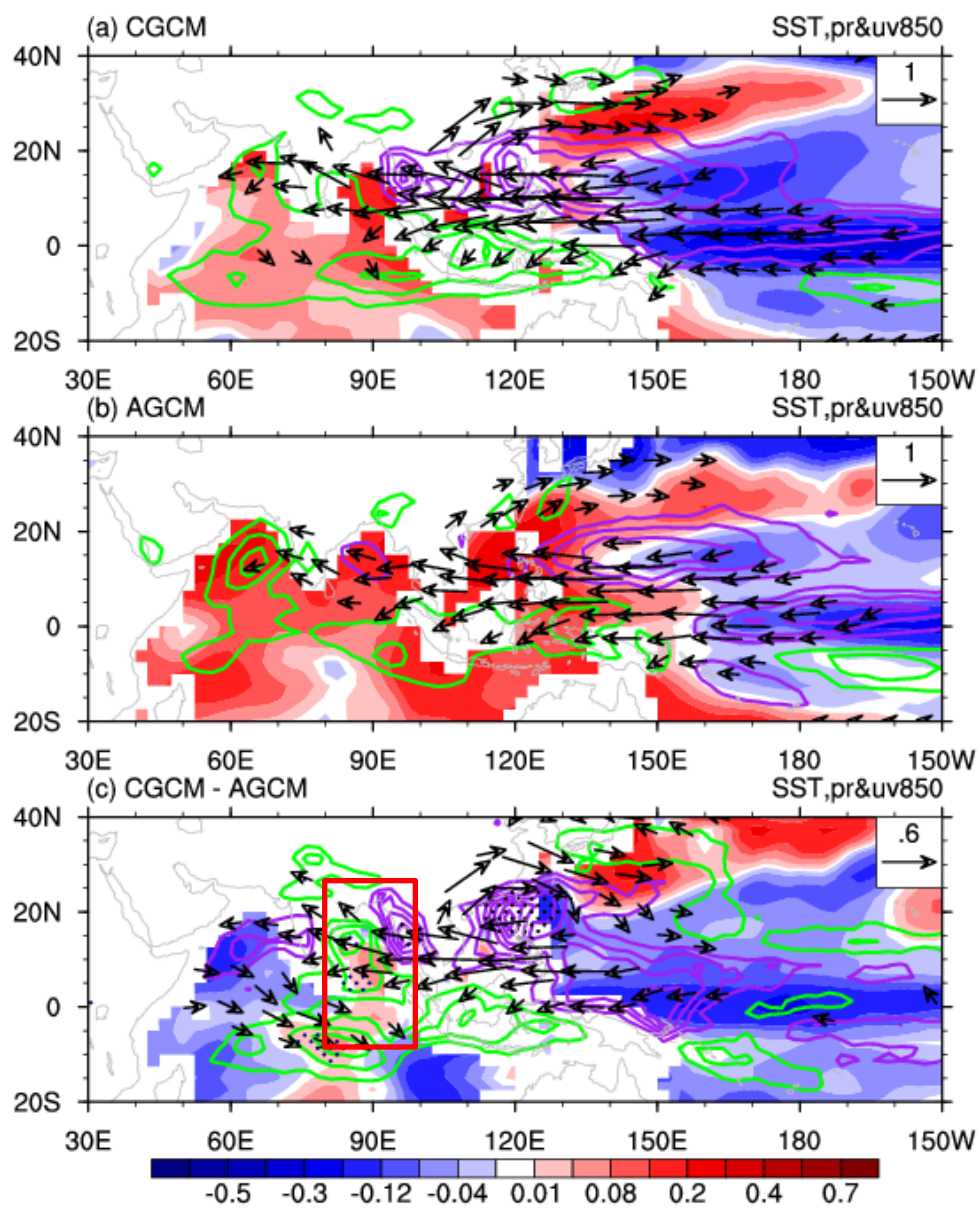
- Better Indian ocean positive precp, better Kelvin wave response.
- CMIP5 MME better than CMIP3 MME



- **17 CMIP5 AGCMs and corresponding CGCMs** are analyzed
- **Observational and reanalysis data:**
 - **NCEP2&ERA40; GPCP&CMAP; ERSST**
- **the period for the comparison between AGCMs and CGCMs is 1979-2005**
- **All the datasets are interpolated into common grid 2.5°x2.5°**



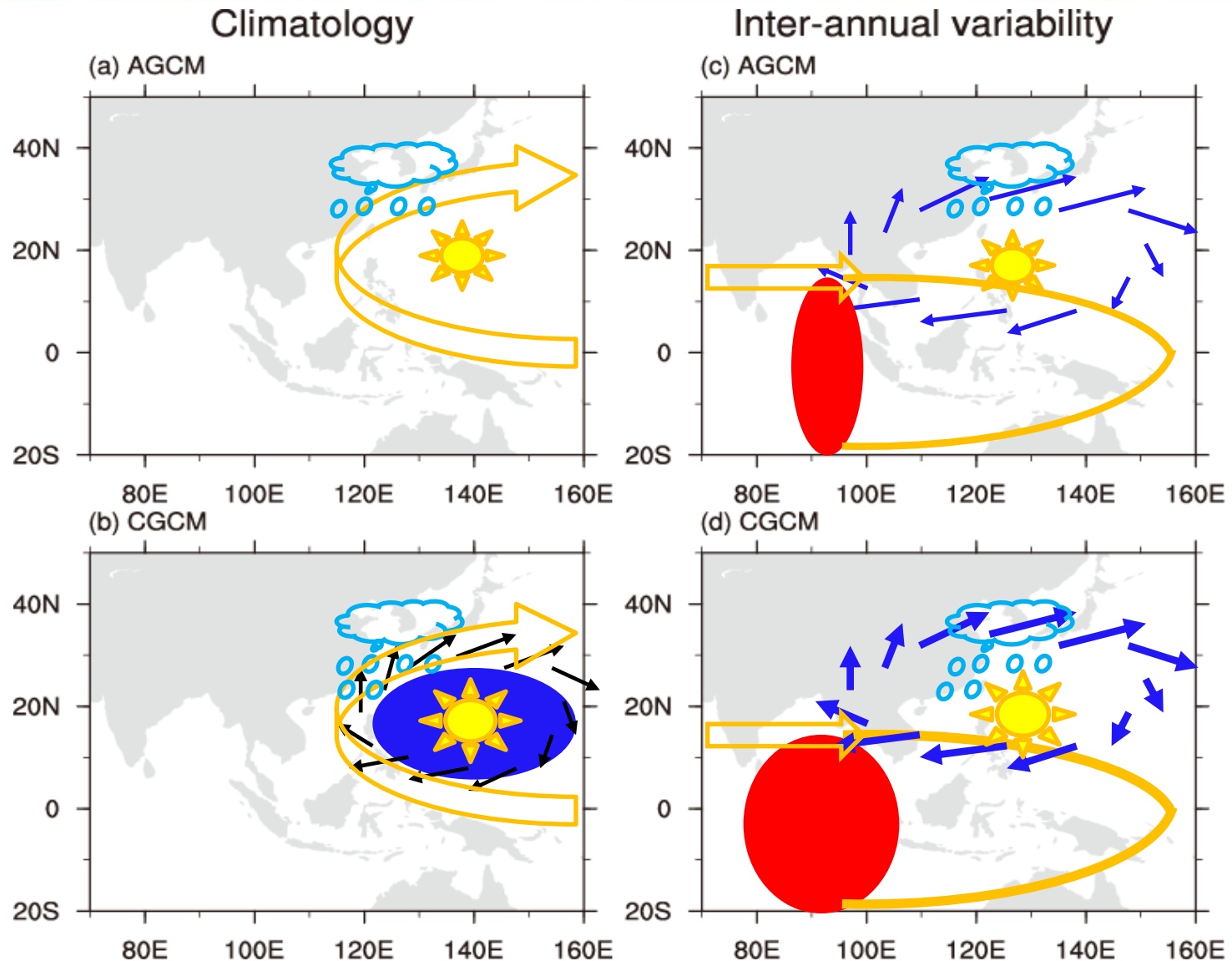
Anomalies of SST, precipitation, and 850 hPa winds in El Nino decaying year summer



Shading: SST
Green contour: positive precipitation
Purple contour: negative precipitation
Vector: 850 hPa winds

- ◆ CGCM: SSTA over TEIO is warmer than the OBS.
- ◆ **Warmer TEIO SSTA** -> more precipitation -> stronger Kelvin wave response as W. Pac AC -> **enhanced EASM simulation.**
- ◆ Local colder SST over the W. Pac also enhances the W. Pac AC

Schematic plot of the air-sea coupling's role in the EASM simulation



Point # 2



◆ Biases of AGCM:

Northward shift of the WP subtropical high in mean state;

Southward shift of the WP AC in interannual variability.

◆ Improvements of CGCM

Mean state: Better WPSH at a cost of colder local SST.

Interannual variability: Improvements in WP AC location and intensity of monsoon rainfall anomaly, due to the enhanced IO-WPAC teleconnection through the air-sea coupling.

◆ Dynamics:

More rainfall over the Indian Ocean associated with a warmer SST, and a stronger equatorial Kelvin wave response in the W. Pacific.



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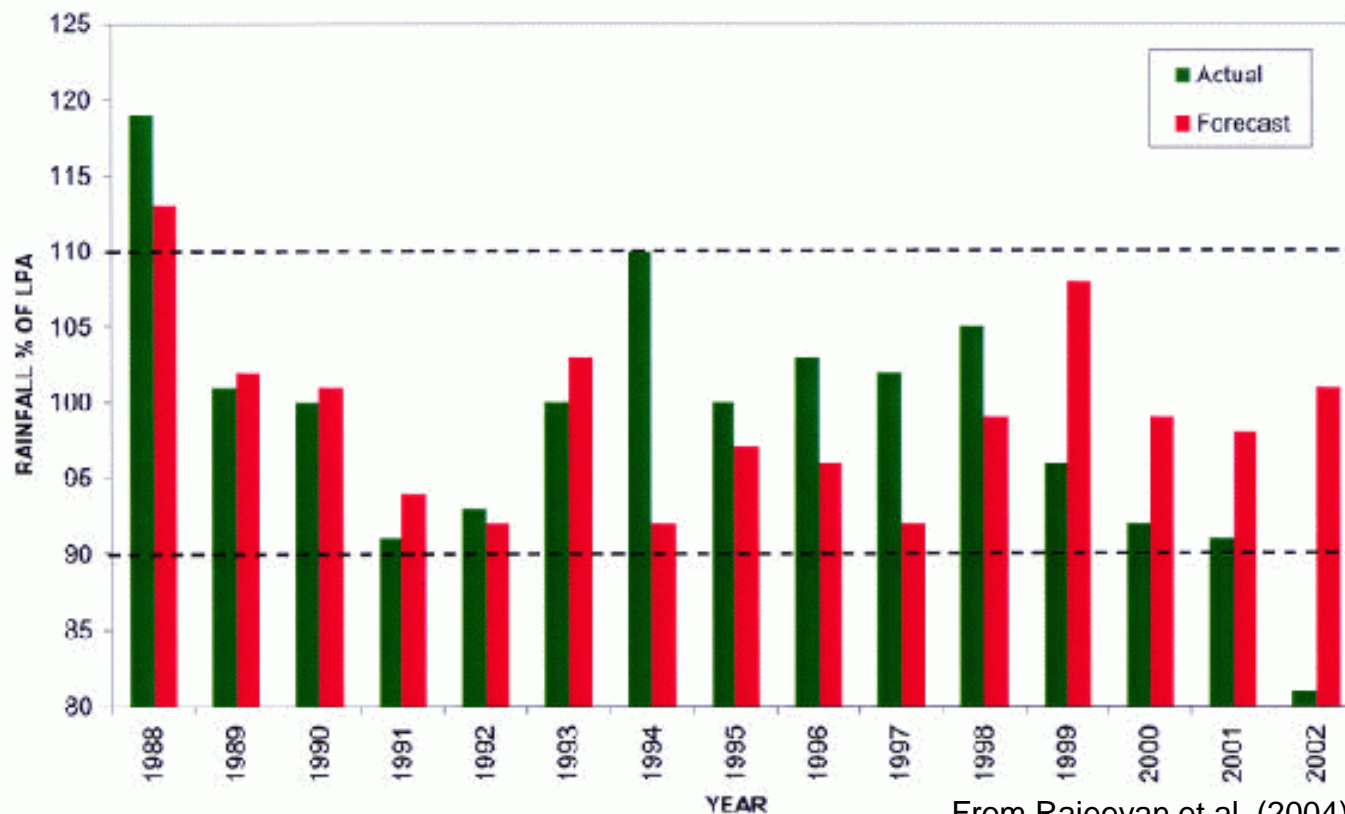
(V. Old) Indian monsoon statistical forecasting

The 10 parameters (and their correlation coefficients with AIR*) are:

1. **Arabian Sea SST** (Jan and Feb) 0.55
2. **Eurasian snow cover** (Dec) -0.46
3. **NW Europe Temperature** (Jan) 0.46
4. **NINO3 SST anomaly** (Jul-Sep previous year) 0.42
5. **South Indian Ocean SST** (Mar) 0.47
6. **East Asia Pressure** (Feb and Mar) 0.61
7. **Northern Hemisphere 50 hPa wind pattern** (Jan) -0.51
8. **Europe Pressure Gradient** (Jan) 0.42
9. **South Indian Ocean 850 hPa zonal wind** (Jun) -0.45
10. **NINO3.4 SST tendency** (between Jan and Jun) -0.46

*AIR = All India Rainfall

Statistical forecast performance



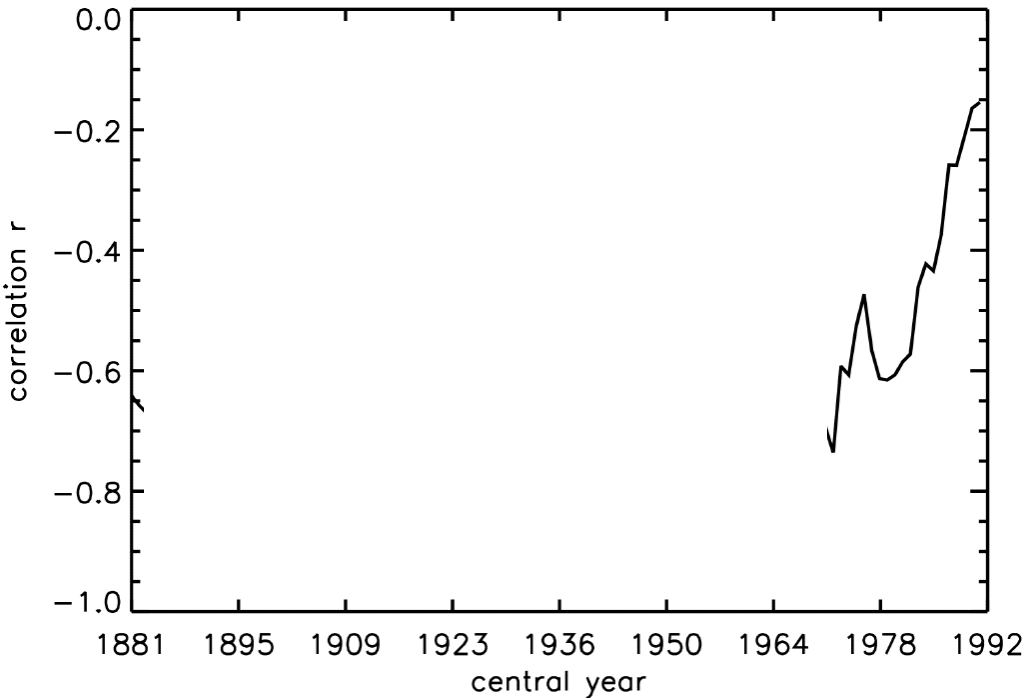
From Rajeevan et al. (2004) *Current Science*

- ❖ Performance of the previous IMD model (16 parameter power regression)
- ❖ Note the gradual deterioration in skill and the failure to predict the 2002 drought
- ❖ The correlations between *predictors* and *predictands* are not necessarily stationary in time, so [dynamical models](#) (coupled ocean-atmosphere GCMs) are beginning to be used for seasonal forecasting of the monsoon

Courtesy: Andy Turner



Moving correlation between AIR and Niño-3 SST during JJAS



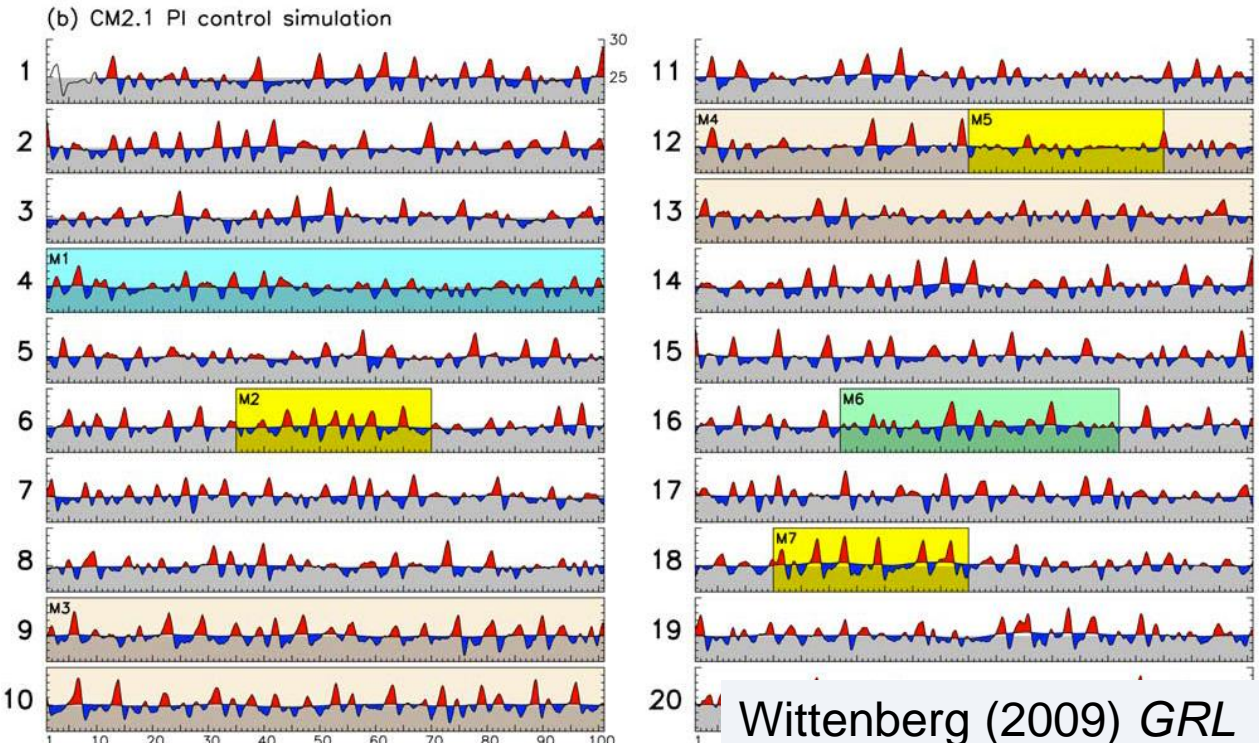
- ❖ The monsoon-ENSO teleconnection has been characterized by apparent recent weakening, but...
- ❖ Considerable interdecadal variability in the past
- ❖ Recent El Niño events (2002, 2004, 2009) have again been related to monsoon droughts of (81%, 87%, 78% LPA AIR)

Is recent “weakening” related to warming (e.g. Krishna Kumar *et al.*, 1999)?



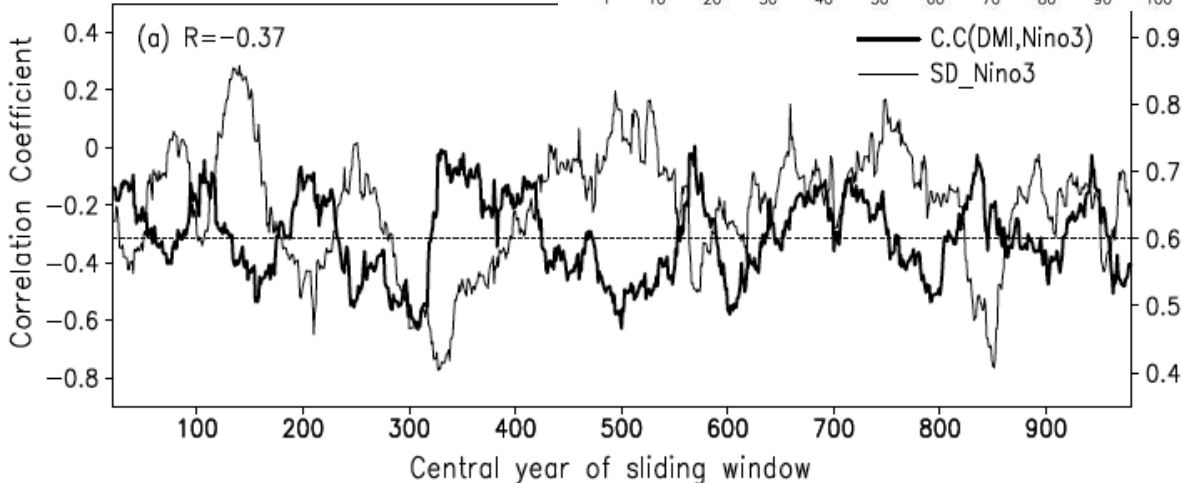
ENSO variance and variations in the monsoon-ENSO teleconnection

- ❖ Ability of ENSO to vary internally
- ❖ Modulation of ENSO variance can alter teleconnection



Wittenberg (2009) *GRL*

Chen et al. (2010) *GRL*



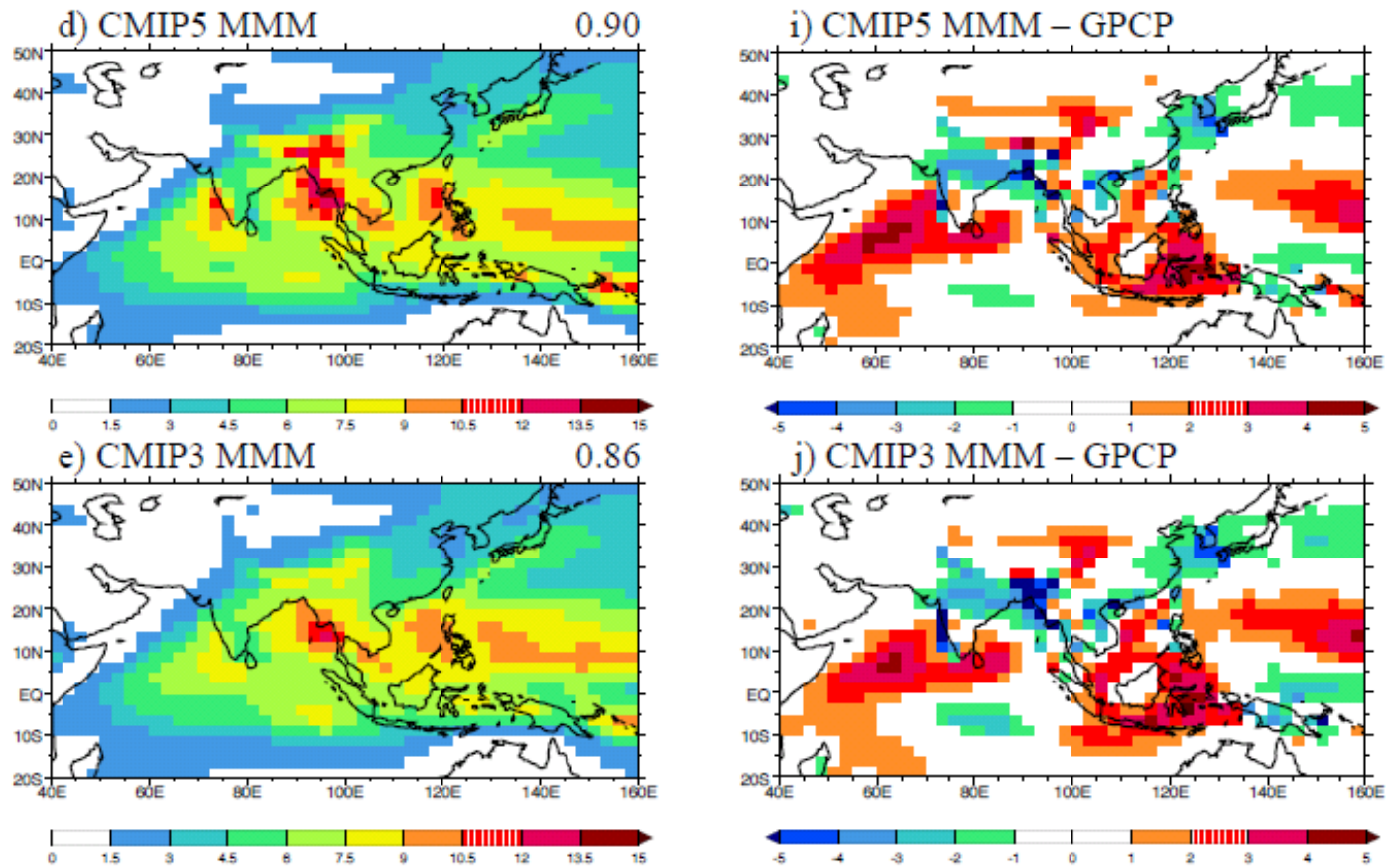
← remember: negative correlation=strong mE teleconnection



CMIP5 and operational models' performances?



Mean JJAS precipitation (left) and bias versus GPCP obs (right)

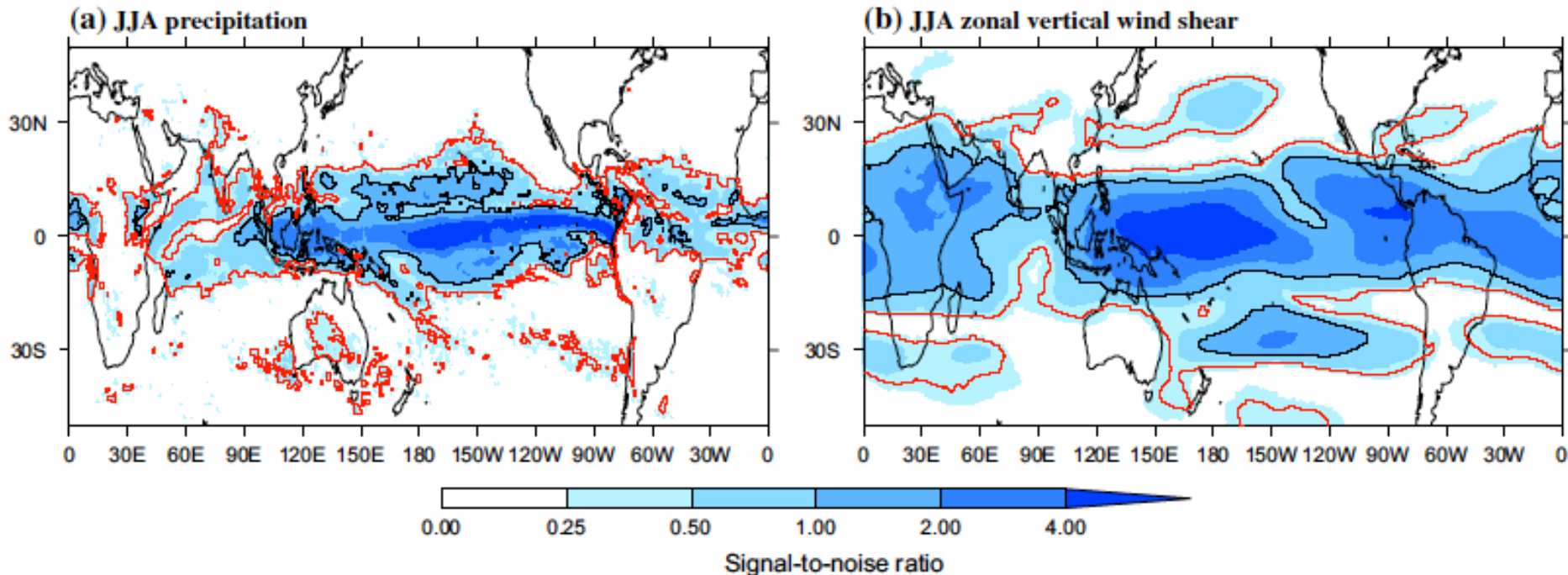


Large biases in CMIP3 and CMIP5 models

Performance in the MetUM GloSea5



MetUM shows more signal in Asian monsoon region for circulation



S/N defined as ratio of variance of interannual timeseries of ensemble mean to time-mean of variances of ensemble for each year

Performance in the MetUM GloSea5



MetUM shows more signal in Asian monsoon region for circulation

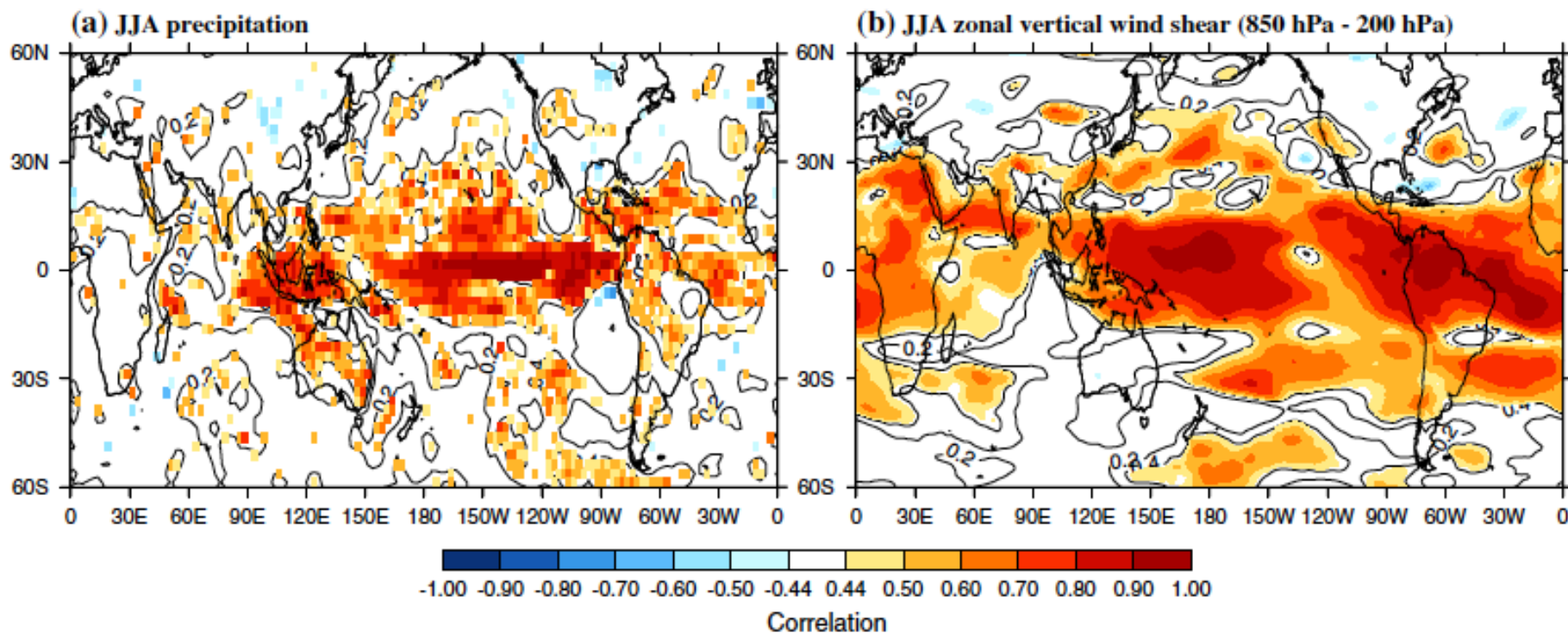
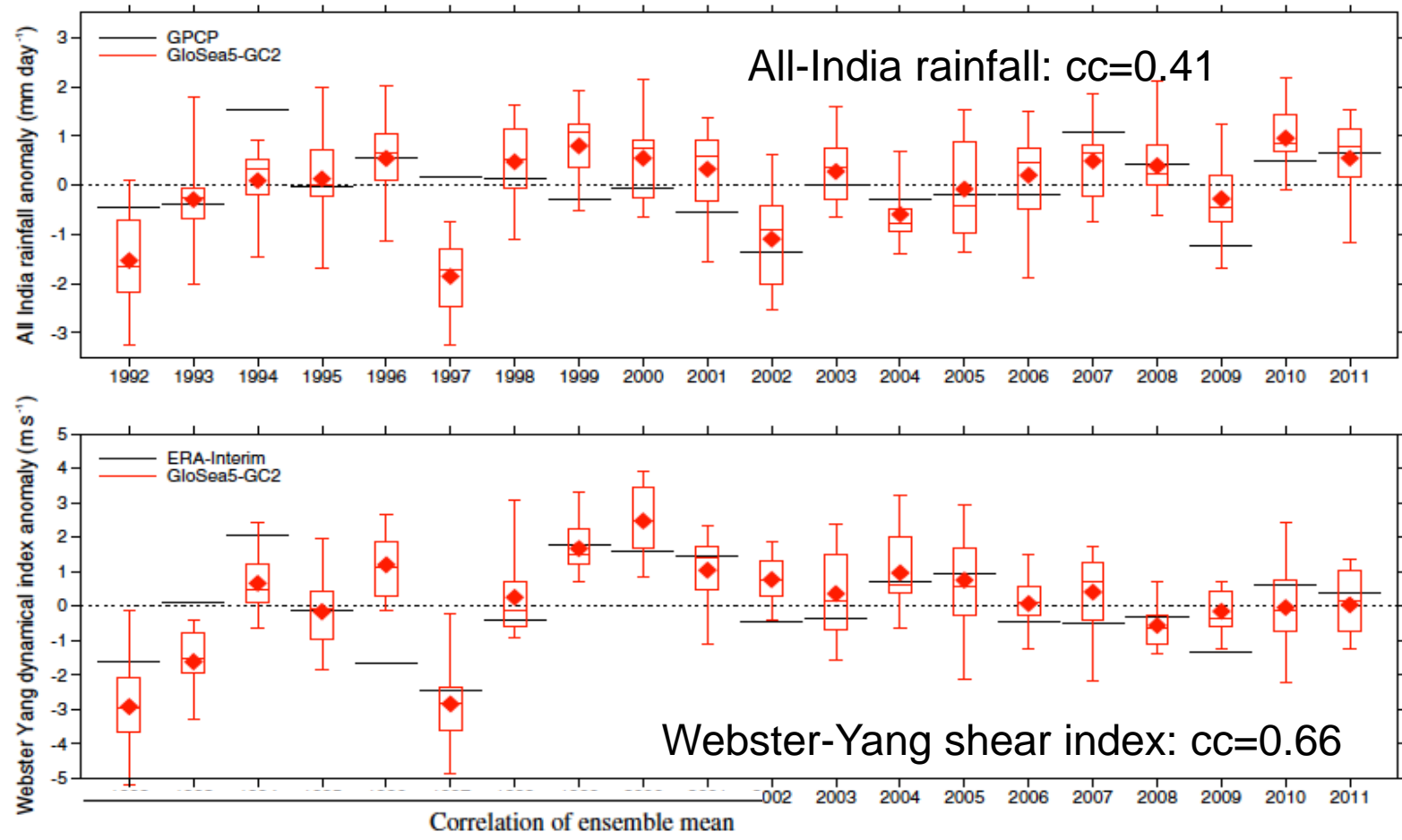


Fig. 3 Grid-point anomaly correlations of GPCP JJA precipitation and ERA-Interim JJA vertical wind shear with their GloSea5-GC2 ensemble mean equivalents. Significant skill (0.44, $p < 0.05$) is shaded, while lower skill is contoured at 0.2 and 0.4



Large-scale circulation measures outperform localized rainfall



AIR	0.41
Wang-Fan index	0.36
Webster-Yang index	0.66

Point # 3



- ◆ The Indian summer monsoon prediction traditionally relies on statistical model, but the recent decades witnessed a gradual deterioration in skill and the failure to predict the 2002 drought.
- ◆ The monsoon-ENSO teleconnection has been characterized by apparent recent weakening. Modulation of ENSO variance can alter the monsoon-ENSO teleconnection.
- ◆ CMIP models show large biases in monsoon rainfall simulation. There exists intimate connection between biases in monsoon circulation and precipitation.
- ◆ MetUM GloSea-5 shows more signal in Asian monsoon region for circulation.



THANKS

<http://www.lasg.ac.cn/staff/ztj>